## Spherical shell coordinates

```
clear
set_demo_defaults;
R_märs = 3389508; % [m] Mars' mean radius
grav = 3.711; % [m/s^2] grav. acceleration on Mars
```

We have seen that moving to cylindrical coordinates removed the ambiguity in the interpretation of the polar recharge. To properly incorporate precipitation we need to go to a geometry with a meaningful surface area. In 1D linear coordinates the surface area is an arbitrary function of the undetermined width. In cylindrical coordinates we have a proper surface area, but given that the southern highlands aquifer stretches halfway through the northern hemisphere the we have a huge error in the actual surface area compared to a sphere, see figure. The same would be true for any estimate of the actual groundwater volume.

```
theta_bnd = acos(1/3);
theta_bnd_deg = rad2deg(theta_bnd)
theta_bnd_deg = 70.5288
```

```
theta_vec = linspace(0,90+theta_bnd_deg,100);
l = R_mars*deg2rad(theta_vec); % [m] distance to dichotomy bnd
A_cyl = pi*l.^2;
A_cap = 2*pi*R_mars^2*(1-cos(deg2rad(theta_vec)));
figure('position',[10 10 900 600])
subplot 121
plot(theta_vec,A_cyl/1e6,theta_vec,A_cap/1e6)
xlabel '0 [\circ]', ylabel 'A [km^2]', pbaspect([1 . 8 1])
legend('Cylinder','Sphere','location','northwest')
subplot }12
plot(theta_vec,(A_cyl-A_cap)./A_cap*100)
xlabel '0 [\\overline{circ]', ylabel''error [%]', pbaspect([1 . 8 1])}
```




## Spherical coordinates

This motivates us to discretize the discrete operators in spherical coordinates. The definition of standard variables in spherical coordinates is shown in the figure below.


Here $r$ is the radial coordinate, $\theta$ is the co-lattitude and $\varphi$ is the circumfrencial coordinate. The associated definition of the gradient and the divergence

$$
\nabla h=\frac{\partial h}{\partial r} \widehat{\boldsymbol{\rho}}+\frac{1}{r} \frac{\partial h}{\partial \theta} \widehat{\boldsymbol{\theta}}+\frac{1}{r \sin \theta} \frac{\partial h}{\partial \varphi} \widehat{\boldsymbol{\rho}}
$$

$\nabla \cdot \mathbf{q}=\frac{1}{r^{2}} \frac{\partial\left(r^{2} q_{r}\right)}{\partial r}+\frac{1}{r \sin (\theta)} \frac{\partial}{\partial \theta}\left(\sin (\theta) q_{\theta}\right)+\frac{1}{r \sin (\theta)} \frac{\partial q_{\varphi}}{\partial \varphi}$
where $\mathbf{q}=\left[q_{r} q_{\theta}, q_{\varphi}\right]$.
The southern highlands aquifer is in a spherical shell, so that $r=R_{\text {Mars }}$ is fixed. To obtain a one-dimensional model we assume no change in the circumferential dirrection, $\varphi$, so that $\partial / \partial \varphi=0$. Therefore, the remaining independent variable is the co-lattitude, $\theta$, and one-dimensional operators in sherical shell geometry, $x=\theta$, are

- $\nabla h=\frac{1}{R_{\text {Mars }}} \frac{\mathrm{d} h}{\mathrm{~d} x}$
- $\nabla \cdot \mathbf{q}=\frac{1}{R_{\text {Mars }} \sin (x)} \frac{\mathrm{d}}{\mathrm{d} x}(\sin (x) q)$

In spherical shell coordinates both the divergence and the gradient change. Again we have to amend the function build_ops.m.

## Discrete operators

The discrete divergence and gradient matrix in spherical shell geometry can therefore be obtained as follows:

```
Grid.xmin = 0.1; Grid.xmax = 1; Grid.Nx = 35;
Grid = build_grid(Grid);
[D,G,I] = build_ops(Grid);
% Modification for spherical shell
Grid.R_shell = R_mars;
Rf = spdiags(sin(Grid.xf),0,Grid.Nx+1,Grid.Nx+1);
Rcinv = spdiags(1./(Grid.R_shell*sin(Grid.xc)),0,Grid.Nx,Grid.Nx);
D = Rcinv*D*Rf;
G = G/Grid.R_shell;
L = -D*G;
```

Similar to the cylindrical coordinates we evaluate terms outside the divergence at cell centers and terms inside the divergence at cell faces.

## Spherical shell aquifer with precipitation

## Dimensional

The equations for the steady confined aquifer with precipitation on a spherical shell are given by
$\frac{1}{R_{\text {Mars }} \sin (\theta)} \frac{\mathrm{d}}{\mathrm{d} \theta}\left(\sin (\theta) b K \frac{1}{R_{\text {Mars }}} \frac{\mathrm{d} h}{\mathrm{~d} x}\right)=q_{p}$ on $\theta \in\left[0, \theta_{b}\right]$
with the boundary conditions

$$
\left.\frac{\mathrm{d} h}{\mathrm{~d} \theta}\right|_{0}=0 \text { on } h\left(\theta_{b}\right)=h_{o}
$$

Th parameter values are as before

```
yr2s=60^2*24*365.25; % second per year
rho = 1e3; % [kg/m^3] desity of water
grav = 3.711; % [m/s^2] grav. acceleration on Mars
k = 1e-11; % [m^2] permeability (Hanna & Phillips 2005)
mu = 1e-3; % [Pa s] water viscosity
ho = -500; % [m] sea level
b = 5e3; % [m] aquifer thickness
theta_bnd = pi-acos(1/3); % [rad] angel dichotomy boundray from south pole
% derived values
K = k*rho*grav/mu; % [m/s] hydraulic conductivity
```


## Dimensionless

The angle, $\theta$, in radiants is defined as ratio of the arc length, $s$, to the radius, $R$, of the circle, $\theta=s / R$, and hence dimensionless and of order one. We define the characteristic scale $h_{c}=q_{p} R_{\text {Mars }}^{2} /(b K)$ and the associated dimensionless head $h^{\prime}=\left(h-h_{o}\right) / h_{c}$. This induces the scale $q_{c}=q_{p} R_{\text {Mars }} / b$ for the flux. The dimensionless equations are

$$
-\frac{1}{\sin \theta} \frac{\mathrm{~d}}{\mathrm{~d} \theta}\left(\sin \theta \frac{\mathrm{~d} h^{\prime}}{\mathrm{d} \theta}\right)=1 \text { on } \theta \in\left[0, \theta_{b}\right]
$$

with the boundary conditions
$\left.\frac{\mathrm{d} h^{\prime}}{\mathrm{d} \theta}\right|_{0}=0$ and $h^{\prime}\left(\theta_{b}\right)=0$.
The dimensionless flux is simply $q^{\prime}=-\frac{\mathrm{d} h}{\mathrm{~d} \theta}$. The only dimensionless parameter is therfore the angle of the dichotomy boundary $\theta_{b}$.

## Analytic solutions

The dimensionless analytic solution is given by
$h^{\prime}=\log \left(\frac{\cos \theta+1}{\cos \theta_{b}+1}\right)$ and $q^{\prime}=\frac{1-\cos \theta}{\sin \theta}=\csc \theta-\cot \theta$.

```
hD_ana = @(theta,theta_bnd) log((cos(theta)+1)/(cos(theta_bnd)+1));
qD_ana = @(theta) csc(theta) - cot(theta);
```

The solution is show in the figure below for increasing co-lattidudes of the dichotomy boundary

```
h_scale_plot = 0.3;
theta_bnd_vec = [theta_bnd:-theta_bnd/5:theta_bnd/5];
figure('position',[10 10 900 600])
for i = 1:length(theta_bnd_vec)
```

```
theta = linspace(0,theta_bnd_vec(i),1e2);
subplot (2,2,1)
plot(theta,hD_ana(theta,theta_bnd_vec(i))); hold on
subplot (2,2,3)
plot(theta,qD_ana(theta)); hold on
subplot(2,2,[2 4])
x_h = (1+hD_ana(theta,theta_bnd_vec(i))*h_scale_plot).*sin(theta);
z_h = (1+hD__ana(theta,theta__ond_vec(i))*h_scale_plot).*cos(theta);
plot(x_h,-z_h,'-','linewidth',1.5), hold on
end
subplot (2, 2,1)
ylabel 'h''', xlabel '0'
subplot (2,2,3)
ylabel 'q''', xlabel '0'
subplot(2,2,[2 4])
theta_sphere = linspace(0,pi,5e2);
x_base = sin(theta_sphere);
z_base = cos(theta_sphere);
plot(x_base,z_base,'k','linewidth',1.5)
axis equal
xlim([0 1.5]), ylim(1.5*[-1 1])
```



Whe looking at the picture of the head on the sphere, keep in mind that this is not the groundwater table, we are still looking at a confined aquifer that looks something like the sketch below


## Numerical solution

The construction of the modified divergence and gradient will be integrated into the function build_ops.m and can be activated by a new field in the Grid structure called Grid.geom = 'spherical_shell';.
Note, for the flux computations the vectors Grid.V and Grid.A have to be aupdated appropritely!
Note also that all the other functions (build_bnd.m, solve_lbvp.m, comp_flux.m) do not need to be updated.

```
theta_ana = linspace(0,theta_bnd,1e2);
Grid.xmin = 0;
Grid.xmax = theta_bnd;
Grid.Nx = le1;
Grid.geom = 'spherical_shell';
Grid.R_shell = 1;
Grid = build_grid(Grid);
% Operators
[D,G,I] = build_ops(Grid);
L = -D*G;
fs = ones(Grid.Nx,1);
% Boundary conditions
BC.dof_dir = [Grid.dof_xmax];
BC.dof_f_dir = Grid.dof_f_xmax;
BC.g = hD_ana(Grid.xc(Grid.dof_xmax),theta_bnd);
BC.dof_neu = [];
BC.dof_f_neu = [];
BC.qb = [];
[B,N,fn] = build_bnd(BC,Grid,I);
hD = solve_lbvp(L,fs,B,BC.g,N);
qD = comp_flux(D,1,G,hD,fs,Grid,BC);
```

```
figure('position',[10 10 900 600])
subplot 121
plot(theta_ana,hD_ana(theta_ana,theta_bnd)), hold on
plot(Grid.xc,hD,'o','markerfacecolor','w','markersize', 8)
xlabel '0 [rad]', ylabel 'h'' ', pbaspect([1 . 8 1])
legend('analytic','numeric','location','southwest')
subplot 122
plot(theta_ana,qD_ana(theta_ana)), hold on
plot(Grid.\overline{xf,qD,'o','markerfacecolor','w','markersize', 8)}
xlabel '0 [rad]', ylabel 'q'' ', pbaspect([1 . 8 1])
legend('analytic','numeric','location','northwest')
```



