

# Downward oxidant transport through Europa's ice shell in porosity waves

Marc Hesse<sup>1,2</sup>, Steven Vance<sup>3</sup>, Jacob Jordan<sup>4</sup>

<sup>1</sup>Department of Geological Sciences, University of Texas at Austin

<sup>2</sup>Institute for Computational Engineering and Sciences, University of Texas at Austin

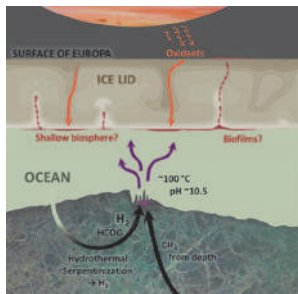
<sup>3</sup>Jet Propulsion Laboratory, California Institute of Technology

<sup>4</sup>Department of Geology & Geophysics, Yale University



January 22, 2019

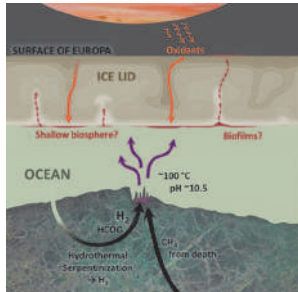
# Redox disequilibria as energy source for life



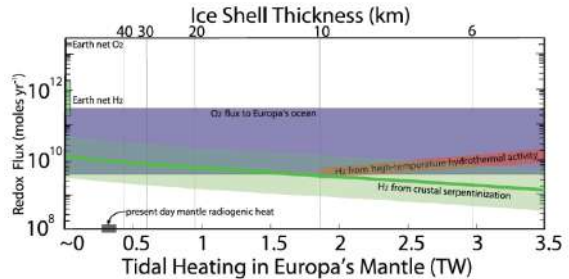
Russell et al. (2017)

- Requires downward oxidant transport through the ice-shell.

# Redox disequilibria as energy source for life



Russell et al. (2017)

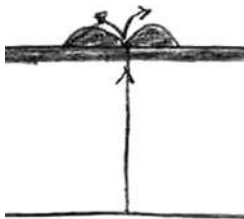


Vance et al. (2016)

- Requires downward oxidant transport through the ice-shell.
- What are the physics of the transfer processes?

# Proposed ice-shell transfer processes

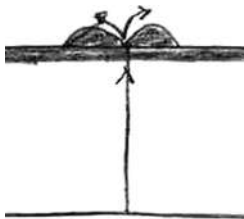
## Resurfacing



Greenberg (2010)

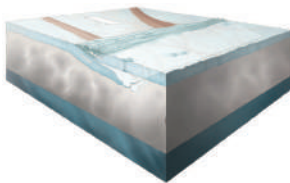
# Proposed ice-shell transfer processes

## Resurfacing



Greenberg (2010)

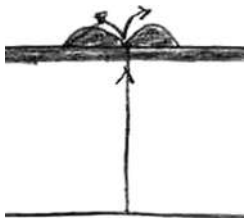
## Subduction



Kattenhorn and  
Prockter (2014)

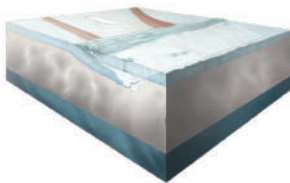
# Proposed ice-shell transfer processes

## Resurfacing



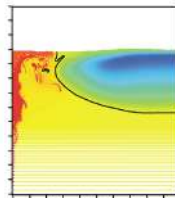
Greenberg (2010)

## Subduction



Kattenhorn and  
Prockter (2014)

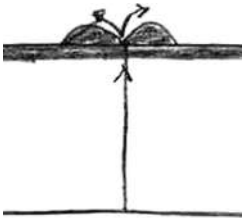
## Impact



Cox and Bauer  
(2015)

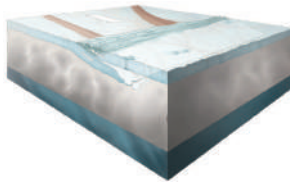
# Proposed ice-shell transfer processes

## Resurfacing



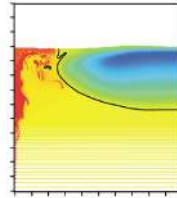
Greenberg (2010)

## Subduction



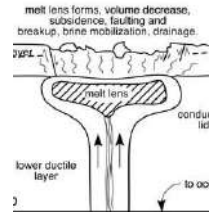
Kattenhorn and  
Prockter (2014)

## Impact



Cox and Bauer  
(2015)

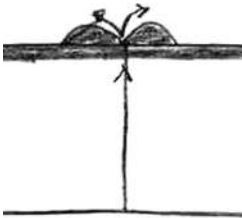
## Brine percolation



Sotin et al. (2002)

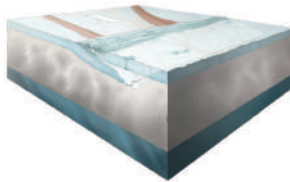
# Proposed ice-shell transfer processes

## Resurfacing



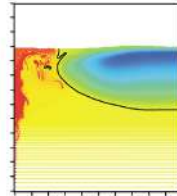
Greenberg (2010)

## Subduction



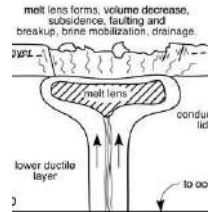
Kattenhorn and  
Prockter (2014)

## Impact



Cox and Bauer  
(2015)

## Brine percolation



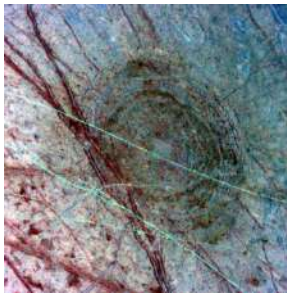
Sotin et al. (2002)

Here we focus on oxidant transport by downward brine percolation.



# Surface features indicating near surface brines

## Impact craters



NASA

## Chaos terrains



NASA

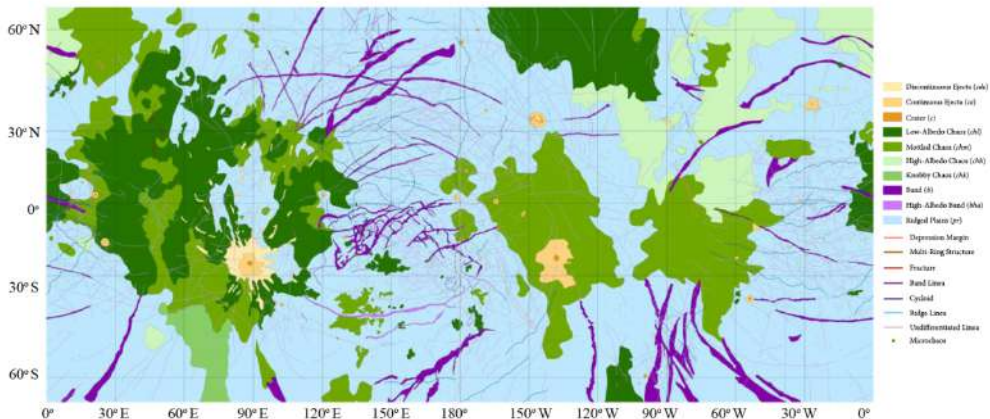
## Lenticulae (domes)



NASA

Assume near-surface brines form in region saturated with oxidants.

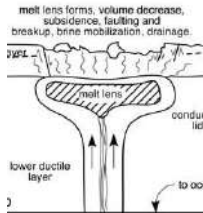
# Distribution of chaotic terrains



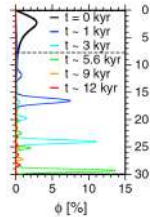
Senske et al. (2018)

# Does brine percolate downward or not?

## Efficient downward percolation



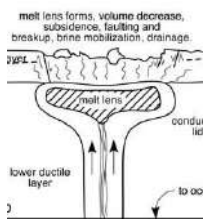
Sotin (2002)



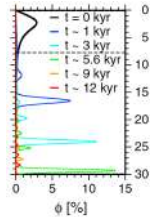
Kalousová (2014)

# Does brine percolate downward or not?

## Efficient downward percolation

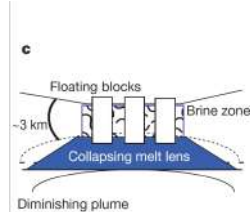


Sotin (2002)

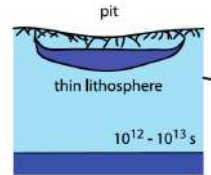


Kalousová (2014)

## Formation of perched aquifers



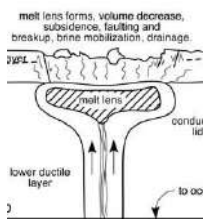
Schmidt (2011)



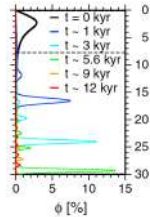
Manga (2017)

# Does brine percolate downward or not?

## Efficient downward percolation

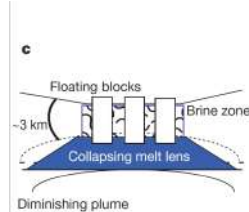


Sotin (2002)

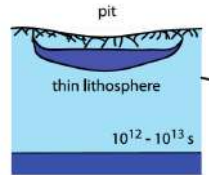


Kalousová (2014)

## Formation of perched aquifers



Schmidt (2011)

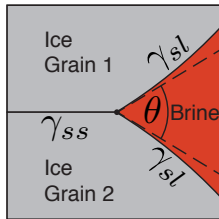


Manga (2017)

Type of behavior is determined by permeability of underlying crust.

- Are small amounts of partial melt present throughout the crust?
- Does this partial melt form a connected network?

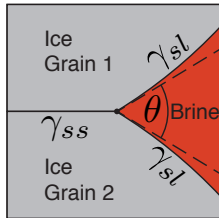
## Dihedral angle $\theta$



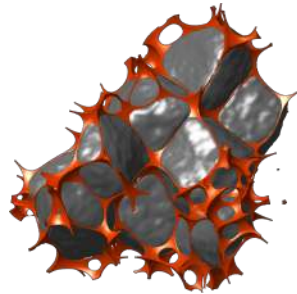
$$\frac{\gamma_{ss}}{\gamma_{sl}} = \frac{\cos(\theta/2)}{2}$$

# Equilibrium melt percolation

Dihedral angle  $\theta$



Wetting:  $\theta \leq 60^\circ$

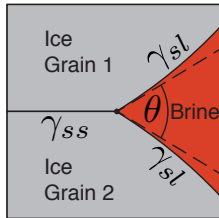


$$\frac{\gamma_{ss}}{\gamma_{sl}} = \frac{\cos(\theta/2)}{2}$$

Percolation

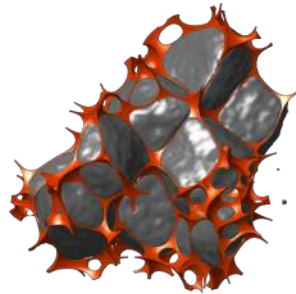
# Equilibrium melt percolation

Dihedral angle  $\theta$



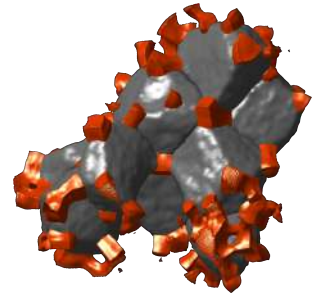
$$\frac{\gamma_{ss}}{\gamma_{sl}} = \frac{\cos(\theta/2)}{2}$$

Wetting:  $\theta \leq 60^\circ$



Percolation

Non-wetting:  $\theta > 60^\circ$

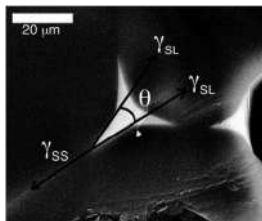


Disconnected



# Does brine connect?

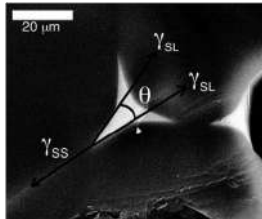
Ice-brine dihedral angle



McCarthy (2012)

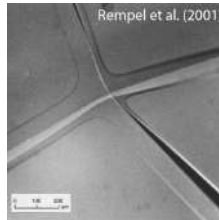
# Does brine connect?

Ice-brine dihedral angle



McCarthy (2012)

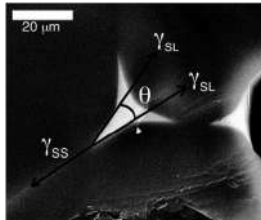
Brine pore network



Rempel (2001)

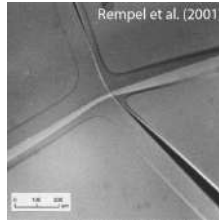
# Does brine connect?

Ice-brine dihedral angle



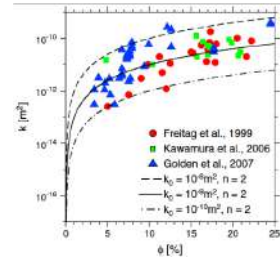
McCarthy (2012)

Brine pore network



Rempel (2001)

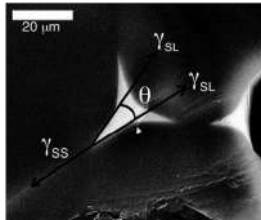
Permeability of ice



Kalousová (2014)

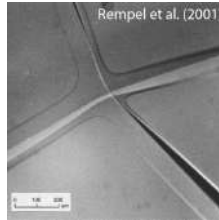
# Does brine connect?

Ice-brine dihedral angle



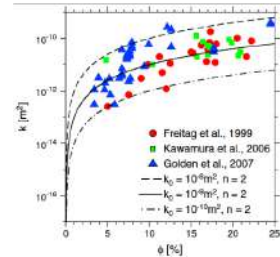
McCarthy (2012)

Brine pore network



Rempel (2001)

Permeability of ice



Kalousová (2014)

In a partially molten ice shell brine is mobile at low melt fraction.

## Mass conservation:

$$\text{Brine:} \quad (\rho_b \phi)_t + \nabla \cdot (\phi \rho_b \mathbf{v}_b) = 0$$

$$\text{Ice:} \quad (\rho_i (1 - \phi))_t + \nabla \cdot ((1 - \phi) \rho_i \mathbf{v}_i) = 0$$

where  $\phi$  is the porosity,  $\rho_p$  and  $\mathbf{v}_p$  are density and velocity of the  $p$ -phase.

## Mass conservation:

$$\text{Brine:} \quad (\rho_b \phi)_t + \nabla \cdot (\phi \rho_b \mathbf{v}_b) = 0$$

$$\text{Ice:} \quad (\rho_i (1 - \phi))_t + \nabla \cdot ((1 - \phi) \rho_i \mathbf{v}_i) = 0$$

where  $\phi$  is the porosity,  $\rho_p$  and  $\mathbf{v}_p$  are density and velocity of the  $p$ -phase.

## Constitutive laws:

$$\text{Compaction relation:} \quad p = p_b - p_i = \xi_\phi \nabla \cdot \mathbf{v}_i$$

$$\text{Darcy's law:} \quad \mathbf{q} = \phi(\mathbf{v}_b - \mathbf{v}_i) = -k_\phi / \mu (\nabla p + \Delta \rho g \hat{\mathbf{z}})$$

where  $k_\phi$  is permeability and  $\xi_\phi$  is bulk viscosity and  $\Delta \rho$  is density difference

## Mass conservation:

Brine:  $(\rho_b \phi)_t + \nabla \cdot (\phi \rho_b \mathbf{v}_b) = 0$

Ice:  $(\rho_i(1 - \phi))_t + \nabla \cdot ((1 - \phi)\rho_i \mathbf{v}_i) = 0$

where  $\phi$  is the porosity,  $\rho_p$  and  $\mathbf{v}_p$  are density and velocity of the  $p$ -phase.

## Constitutive laws:

Compaction relation:  $p = p_b - p_i = \xi_\phi \nabla \cdot \mathbf{v}_i$

Darcy's law:  $\mathbf{q} = \phi(\mathbf{v}_b - \mathbf{v}_i) = -k_\phi/\mu (\nabla p + \Delta\rho g \hat{\mathbf{z}})$

where  $k_\phi$  is permeability and  $\xi_\phi$  is bulk viscosity and  $\Delta\rho$  is density difference

## Governing equations for $\phi$ and $p$ are

Porosity evolution:  $\phi_t + \nabla \cdot [\phi \mathbf{v}_i] = \nabla \cdot \mathbf{v}_i$

Two-phase continuity:  $-\nabla \cdot [\mathbf{q} + \mathbf{v}_i] = 0$

## Mass conservation:

$$\text{Brine:} \quad (\rho_b \phi)_t + \nabla \cdot (\phi \rho_b \mathbf{v}_b) = 0$$

$$\text{Ice:} \quad (\rho_i (1 - \phi))_t + \nabla \cdot ((1 - \phi) \rho_i \mathbf{v}_i) = 0$$

where  $\phi$  is the porosity,  $\rho_p$  and  $\mathbf{v}_p$  are density and velocity of the  $p$ -phase.

## Constitutive laws:

$$\text{Compaction relation:} \quad p = p_b - p_i = \xi_\phi \nabla \cdot \mathbf{v}_i$$

$$\text{Darcy's law:} \quad \mathbf{q} = \phi(\mathbf{v}_b - \mathbf{v}_i) = -k_\phi / \mu (\nabla p + \Delta \rho g \hat{\mathbf{z}})$$

where  $k_\phi$  is permeability and  $\xi_\phi$  is bulk viscosity and  $\Delta \rho$  is density difference

## Governing equations for $\phi$ and $p$ are

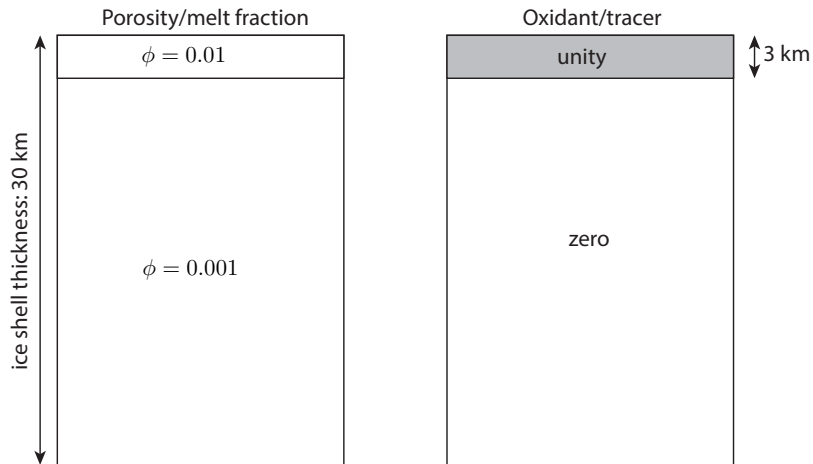
$$\text{Porosity evolution:} \quad \phi_t + \nabla \cdot [\phi \mathbf{v}_i] = \nabla \cdot \mathbf{v}_i$$

$$\text{Two-phase continuity:} \quad -\nabla \cdot [\mathbf{q} + \mathbf{v}_i] = 0$$

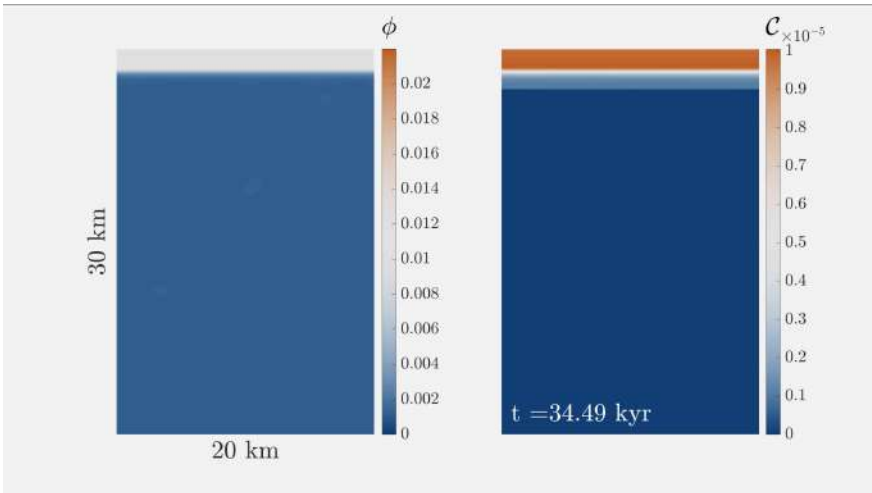
**Internal length-scale:**  $\delta = \sqrt{k_\phi \xi_\phi / \mu}$  (compaction length)



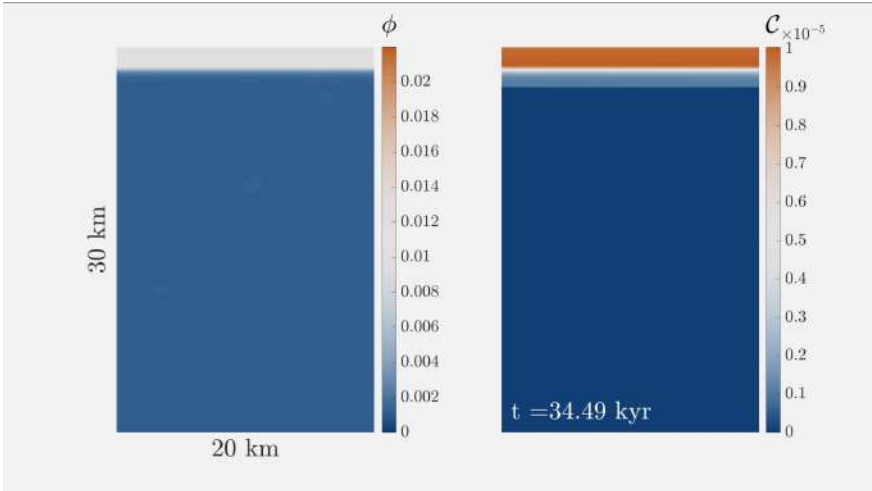
# Simple model problem for oxidant transport



# Slow uniform transport: $\delta/H \sim 1$

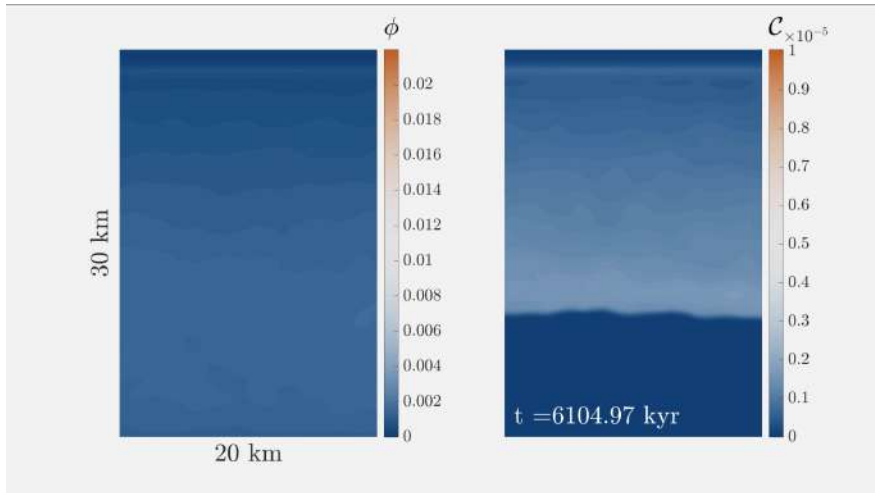


# Slow uniform transport: $\delta/H \sim 1$



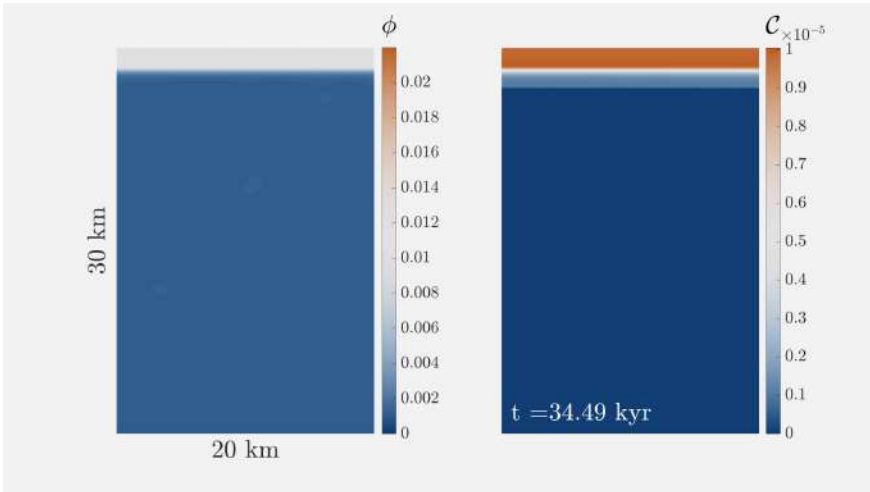
Slow oxidant transport in uniform front as entire crust is dilated.

# Just in case the movie fails!

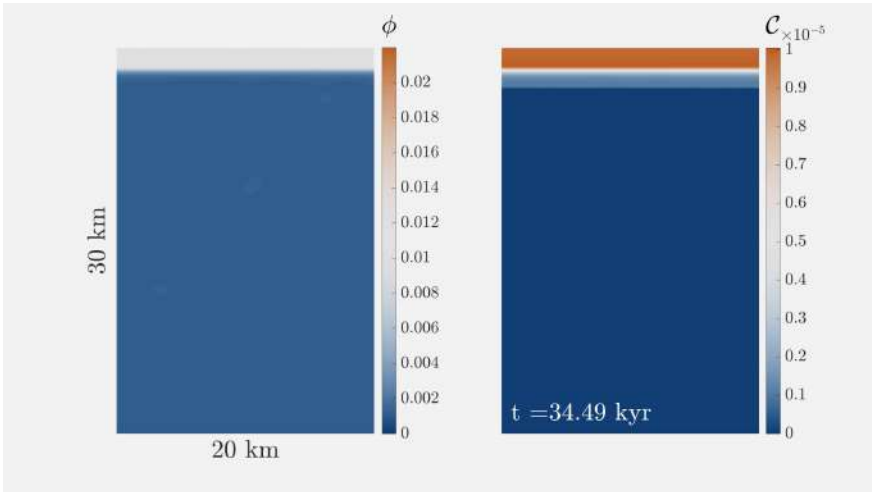


Slow oxidant transport in uniform front as entire crust is dilated.

# Fast localized transport: $\delta/H \ll 1$

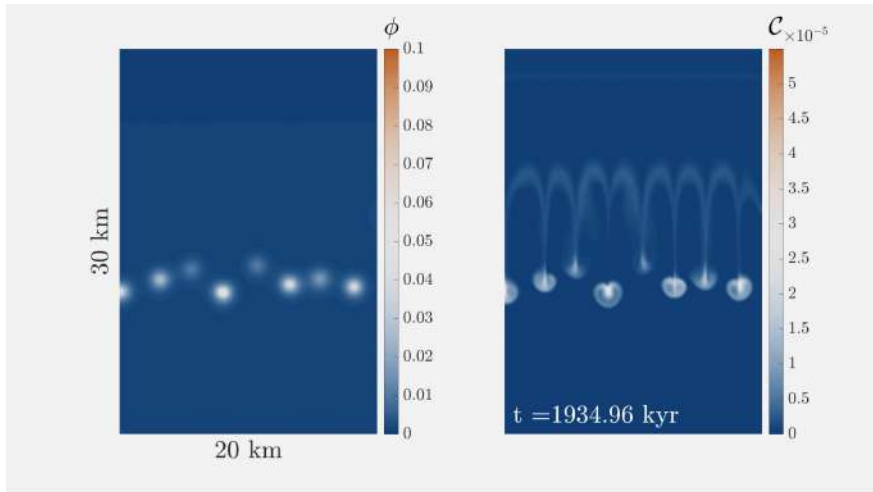


Fast localized transport:  $\delta/H \ll 1$



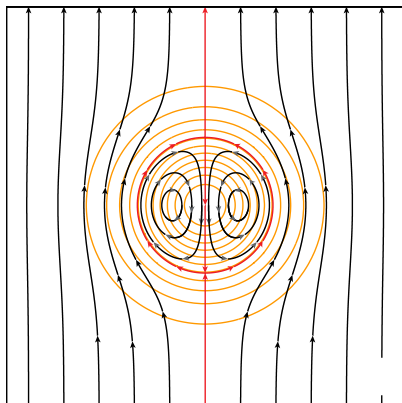
Faster oxidant transport that localizes into 2D porosity waves.

# Just in case the movie fails!



Faster oxidant transport that localizes into 2D porosity waves.

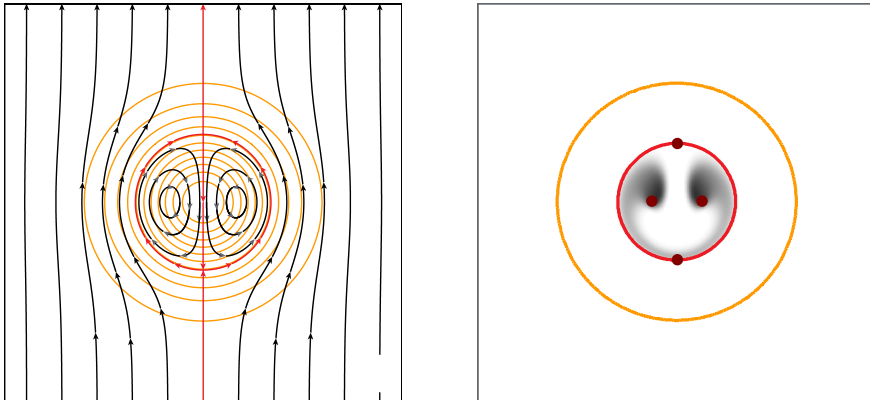
# Fluid recirculation in 2D porosity waves



Jordan and Hesse (2018)



# Fluid recirculation in 2D porosity waves



Jordan and Hesse (2018)

# Changes in transport dynamics over time?

Dynamics are determined by:

$$\frac{\delta}{H} = \frac{\sqrt{k_{\phi}\xi_{\phi}}}{\sqrt{\mu H}}$$

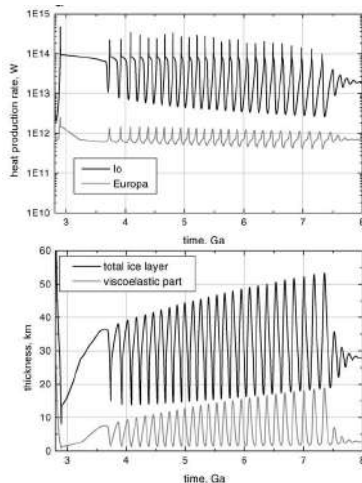
# Changes in transport dynamics over time?

Dynamics are determined by:

$$\frac{\delta}{H} = \frac{\sqrt{k_{\phi}\xi_{\phi}}}{\sqrt{\mu H}}$$

Thermo-orbital evolution leads to

- Large variation in  $H$ .
- Variation in heat production.
  - affects porosity,  $\phi$ .
  - affects permeability,  $k_{\phi}$ .
  - affects bulk viscosity,  $\xi_{\phi}$ .



Husmann and Spohn (2004)

# Changes in transport dynamics over time?

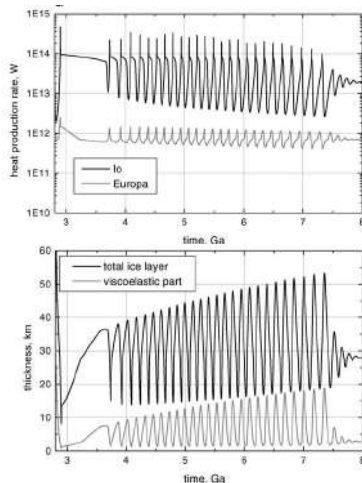
Dynamics are determined by:

$$\frac{\delta}{H} = \frac{\sqrt{k_{\phi}\xi_{\phi}}}{\sqrt{\mu H}}$$

Thermo-orbital evolution leads to

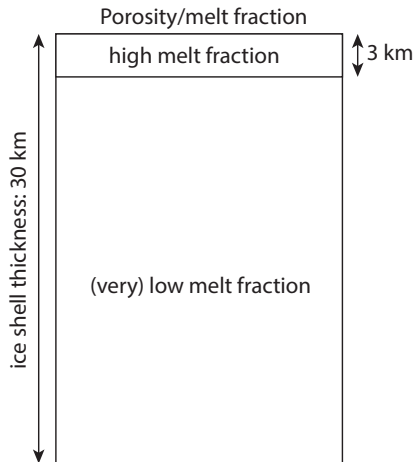
- Large variation in  $H$ .
- Variation in heat production.
  - affects porosity,  $\phi$ .
  - affects permeability,  $k_{\phi}$ .
  - affects bulk viscosity,  $\xi_{\phi}$ .

Explore effect of thermo-orbital evolution on oxidant transport.

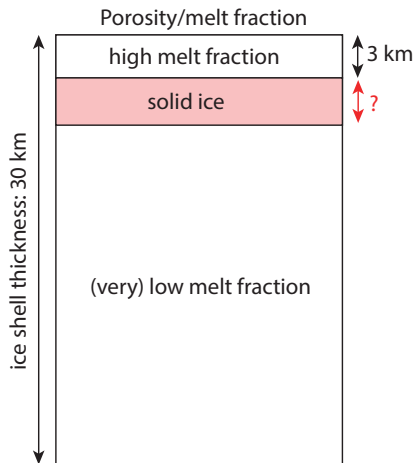


Husmann and Spohn (2004)

# What if there is a layer of solid ice?

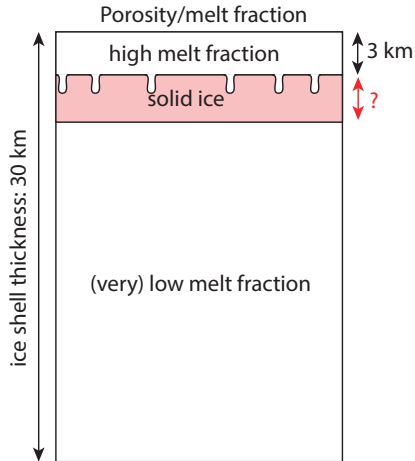


# What if there is a layer of solid ice?



Does solid ice prevent brine drainage?

# What if there is a layer of solid ice?

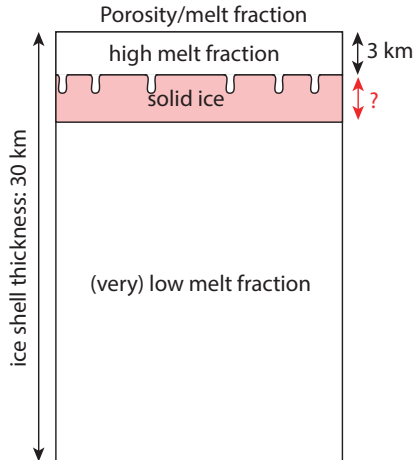


Does solid ice prevent brine drainage?

Possible penetration mechanisms:

- Brine wicks into ice by capillary forces.
- Partial melting induced by latent heat released from brine crystallization.
- Transfer of elastic stresses from the volume expansion of solidifying brine.

# What if there is a layer of solid ice?



Does solid ice prevent brine drainage?

Possible penetration mechanisms:

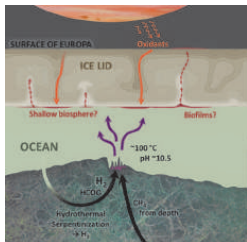
- Brine wicks into ice by capillary forces.
- Partial melting induced by latent heat released from brine crystallization.
- Transfer of elastic stresses from the volume expansion of solidifying brine.

Likely a broad range of behaviors.



# Summary and conclusions

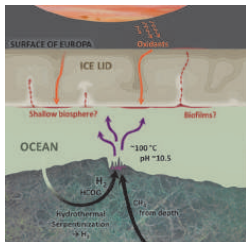
## Redox disequilibria



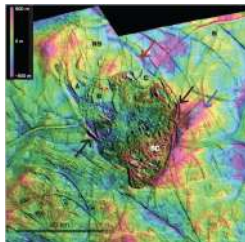
Need O<sub>2</sub> transport.

# Summary and conclusions

## Redox disequilibria



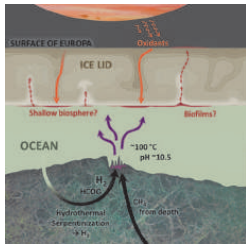
## Near-surface melting



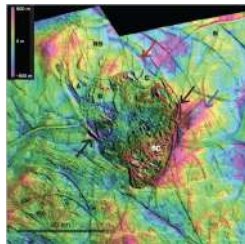
Need O<sub>2</sub> transport. Brine percolation.

# Summary and conclusions

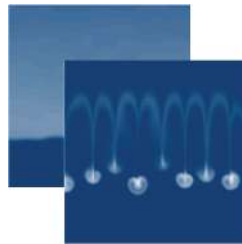
## Redox disequilibria



## Near-surface melting



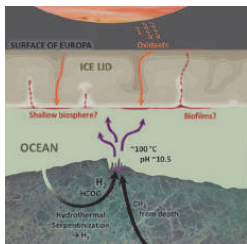
## Transport Regimes



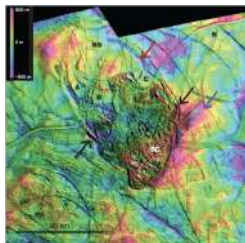
Need  $O_2$  transport. Brine percolation. Change over time.

# Summary and conclusions

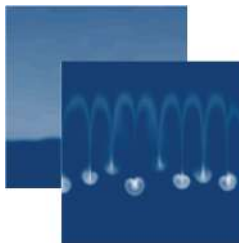
## Redox disequilibria



## Near-surface melting



## Transport Regimes



## To percolate, or not to percolate:



Need O<sub>2</sub> transport. Brine percolation. Change over time. That is the question.

Thank you for your attention.