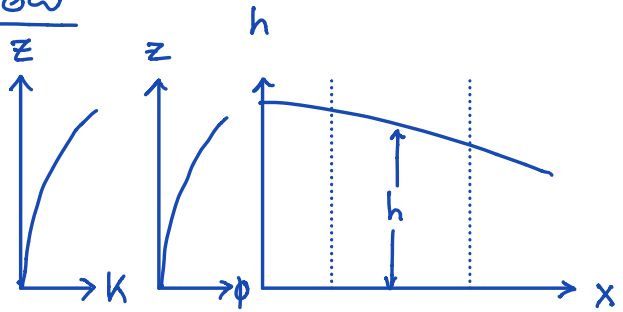


Transient unconfined flow

Let's consider an aquifer with vertical variation of both porosity $\phi(z)$ and hydraulic conductivity $K(z)$.



Any variation can be considered, but for simplicity we consider a power-law variation in both quantities:

$$\phi = \phi_0 z^m \quad K = K_0 z^n \quad \text{typically } \frac{n}{m} \in [2, 3]$$

The balance law for fluid mass is then

$$\frac{\partial}{\partial t} \int_0^h \phi dz + \nabla \cdot \int_0^h q dz = f_s$$

where h is the height of the water table

where the horizontal flux is given by the Dupuit-Forchheimer

approximation: $q = -K \nabla h$

So that we have the following governing equation

for transient unconfined flow

$$\frac{\partial h^{m+1}}{\partial t} - D_h \nabla \cdot [h^{n+1} \nabla h] = f_s$$

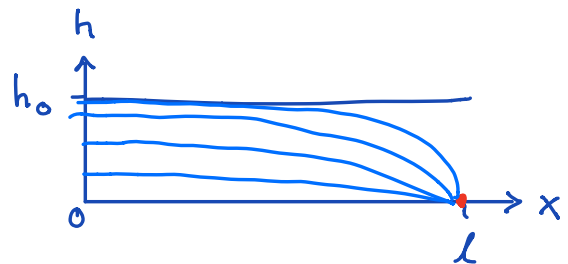
This equation is non-linear both in the flux and in the accumulation term. The two exponents are typically related $n/m \in [2, 3]$. This equation is variably referred to as the Boussinesq eqn or the Porous Media eqn.

Example: Drainage of an unconfined aquifer

$$\text{PDE: } \frac{\partial h^{m+1}}{\partial t} + D_h \nabla \cdot [h^{n+1} \nabla h] = 0$$

$$\text{BC's: } \nabla h \cdot \hat{n}|_0 = 0 \quad h(l) = 0$$

$$\text{IC: } h(x, t=0) = h_0$$



This problem has both an "early" and a "late" self similar soln.

The early soln applies before the head evolution interacts with the boundary at $x=0$.

The late soln applies after the head has started interacting with the bnd at $x=0$ and $h(0, t)$ is dropping.

Here we consider the late solution.

Late self-similar solution

External length scale: $x' = \frac{x}{l}$

External head scale h_0 is not relevant at late time!

⇒ look for internal scale to help find the form of the similarity variable.

assume $h' = \frac{h}{h_c}$ $t' = \frac{t}{t_c}$ substitute into PDE:

$$\frac{h_c^{m+1}}{t_c} \frac{\partial h'^{m+1}}{\partial t'} - D_h \frac{h_c^{n+2}}{l^2} \nabla' \cdot [h'^{n+1} \nabla h'] = 0$$

$$\frac{\partial h'^{m+1}}{\partial t'} - \frac{D_h t_c h_c^{n-m+1}}{l^2} \nabla' \cdot [h'^{n+1} \nabla h'] = 0$$

⇒ this suggests that $h_c^{n-m+1} = \frac{l^2}{D_h t_c}$

This would suffice to non-dimensionalize the equation

but we look for a similarity variable. so we set

$$h = \left(\frac{l^2}{D_h t} \right)^{\frac{1}{n-m+1}} f\left(\frac{x}{l}\right)$$

Substituting this into the PDE results in the self-similar

$$\text{ODE: } \frac{d}{dx'} \left(f^{n+1} \frac{df}{dx'} \right) + \frac{m+1}{n-m+1} f^{m+1} = 0$$

$$\text{BC: } \left. \frac{df}{dx'} \right|_0 = 0 \quad f(1) = 0$$

This non-linear ODE has no known analytic solution and must be solved numerically using the Newton-Raphson Method ∇

However the most important conclusions are gained without even solving the ODE. Consider the mass of GW in the aquifer, given by

$$M(t) = \rho w \int_0^l \int_0^h \phi dz dx = \rho w \int_0^l \frac{\phi_0}{m+1} h^{m+1} dx = \frac{\rho w \phi_0}{m+1} \int_0^l h^{m+1} dz$$

substituting the definition of the similarity variable

$$h = \left(\frac{l^2}{D_h t} \right)^{\frac{1}{n-m+1}} f\left(\frac{x}{l}\right)$$

we have

$$M(t) = \frac{\rho w \phi_0}{m+1} \int_0^1 \left[\left(\frac{l^2}{D_h t} \right)^{\frac{1}{n-m+1}} f(s) \right]^{m+1} l ds$$

$$s = \frac{x}{l}$$

$$dx = l ds$$

so that

$$M(t) = \frac{\rho w \phi_0 l}{m+1} \left(\frac{l^2}{D_h t} \right)^{\frac{m+1}{n-m+1}} \int_0^1 f(s)^{m+1} ds$$

Note: $\int_0^1 f(s)^{m+1} ds$ is just a number ∇

$$M(t) \sim t^{\frac{m+1}{m-n-1}}$$

Example: $m=n=0$ ($k=\text{const.}, \phi=\text{const.}$) $\Rightarrow M \sim t^{-1} = \frac{1}{t}$

$$m=1, n=4 \Rightarrow M \sim t^{\frac{2}{2-4-1}} = t^{-\frac{2}{3}} = \frac{1}{t^{2/3}}$$