

## Lecture 2: Balance Laws

Logistics: Office hrs: Mon 2-3 pm

Wed 3-4 pm

Hydrothermal convection

Last time: - Porous media basics

volume fraction:  $\phi_p = \frac{V_p}{V_T}$

$V_T = \sum V_p$   
porosity

Darcy's law:  $Q = -K A \frac{\Delta h}{\Delta L}$

Rates vs. Fluxes:  $Q$  vs.  $q$

scalar	$\frac{\#}{T}$	vectors	$\frac{\#}{L^2 T}$	discharge	$\frac{L^3}{T}$	spec. disch	$\frac{L^3}{L^2 T} = \frac{L}{T}$
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Today: - Darcy's law in flux form

- General balance law

- Fluid mass balance in porous media

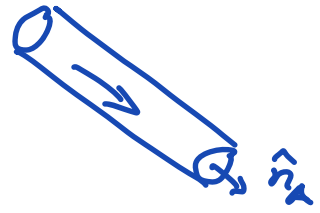
- Incompressible flow in porous media

$\Rightarrow$  first equation to solve

## Vector form of Darcy's law

From Experiments:  $Q = -K A \frac{\Delta h}{\Delta L}$   
↑  
Rate = Discharge

spec. discharge:  $q = \frac{Q}{A} \hat{n}_A$



$$|q| = \frac{Q}{A} = -K \frac{\Delta h}{\Delta L}$$

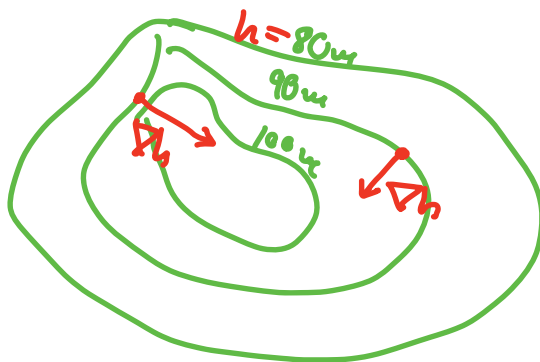
derivative  $\Delta L \rightarrow 0$

$$|q| = -K \frac{dh}{dx} \quad 1D$$

multi dimensional version:

$$q = -K \nabla h$$

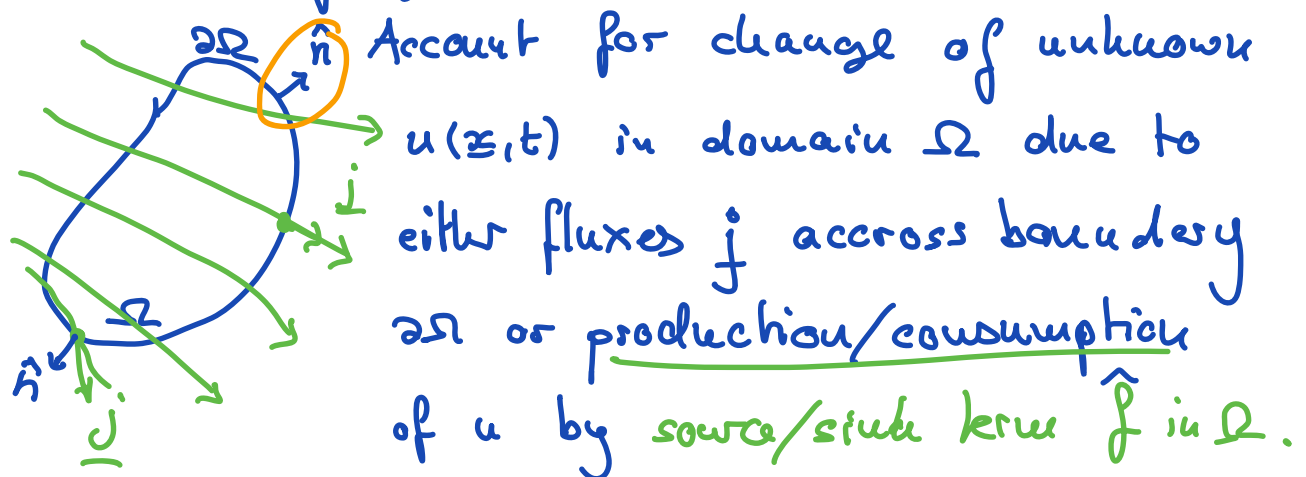
here  $\nabla h = \begin{pmatrix} \frac{\partial h}{\partial x} \\ \frac{\partial h}{\partial y} \\ \frac{\partial h}{\partial z} \end{pmatrix}$



## Balance laws

- most fundamental physical eqns in sciences are balance laws!
- Balance law: accounts for change of some quantity

## Derivation of general balance law



Units of basic quantities:

- $u$  is a density  $\left[ \frac{\#}{L^3} \right]$
- $\underline{j}$  is a flux  $\left[ \frac{\#}{L^2 T} \right]$
- $\hat{f}$  is a vol. source  $\left[ \frac{\#}{L^3 T} \right]$

General balance on  $\Omega$ :

$$\boxed{\frac{dU}{dt} = J + F} \quad \text{macroscopic balance}$$

1)  $U$  is amount of  $u$  in  $\Omega$ :

$$U(t) = \int_{\Omega} u(x,t) dV \quad [\#]$$

2)  $J$  is the rate of transport of  $u$

across  $\partial\Omega$  by  $j$ :  $J(t) = - \int_{\partial\Omega} j(x,t) \cdot \hat{n} \, dA$   ~~$dt$~~

$$[\frac{\#}{T}]$$

3)  $F$  is rate of prod./consumption of  $u$

inside  $\Omega$ :  $F(t) = \int_{\Omega} \hat{f}(x,t) dV$

$$\frac{\#}{T}$$

substitute into balance:

$$\boxed{\frac{d}{dt} \int_{\Omega} u dV = - \int_{\partial\Omega} j \cdot \hat{n} dA + \int_{\Omega} \hat{f} dV} \quad \text{Integral balance law}$$

$$\underbrace{\frac{1}{T} \frac{\#}{L^3} L^3}_{\frac{\#}{T}}$$

$$\underbrace{\frac{\#}{L^2 T} L^2}_{\frac{\#}{T}}$$

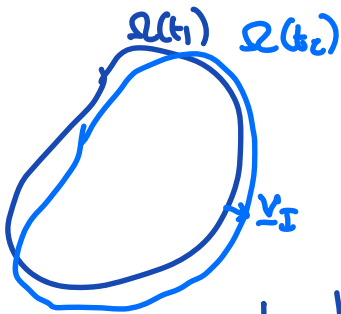
$$\underbrace{\frac{\#}{L^3 T} L^3}_{\frac{\#}{T}}$$

✓

To obtain local PDE we need:

- 1) Exchange the derivative and integral
- 2) Transform surface to a volume integral

### 1) Reynolds Transport Theorem



$$\frac{d}{dt} \int_{\Omega(t)} u(\mathbf{x}, t) dV = \int_{\Omega} \frac{\partial u}{\partial t} dV + \oint_{\partial\Omega} u (\mathbf{v}_I \cdot \hat{\mathbf{n}}) dS$$

In this class we consider Eulerian limit where domain is fixed  $\mathbf{v}_I = 0$

$$\Rightarrow \frac{d}{dt} \int_{\Omega} u dV = \int_{\Omega} \frac{\partial u}{\partial t} dV$$

### 2) Divergence theorem:

$$\int_{\partial\Omega} \mathbf{j} \cdot \hat{\mathbf{n}} dA = \int_{\Omega} \nabla \cdot \mathbf{j} dV$$

Substitute into integral balance law:

$$\int_{\Omega} \underbrace{\left( \frac{\partial u}{\partial t} + \nabla \cdot \mathbf{j} - \hat{\mathbf{f}} \right)}_{g(\mathbf{x}, t)} dV = 0$$

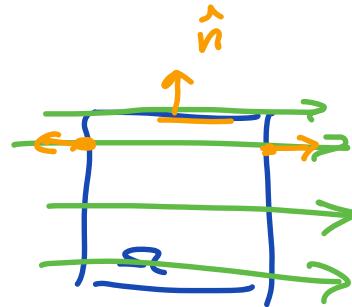
Localization: because  $\Omega$  is arbitrary

the only way for the integral to be always zero is that the integrand itself is zero.

$$\frac{\partial u}{\partial t} + \nabla \cdot \mathbf{j} = \hat{f}$$

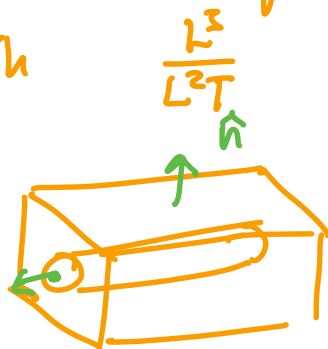
Local form of general balance law

- $u$  = unknown  $[\frac{\#}{L^3}]$
- $\mathbf{j}$  = flux  $[\frac{\#}{L^2 T}]$
- $\hat{f}$  = vol. rat  $[\frac{\#}{L^3 T}]$



Difference between flux & velocity in porous med.

$$\underline{q} = -k \nabla h$$



$$\underline{v} = \frac{\underline{q}}{\phi}$$

## Fluid Mass Balance:

Gen. balance law:  $\frac{\partial u}{\partial t} + \nabla \cdot \mathbf{j} = \hat{f}$

1) Define unknown to be balanced

$$u = \phi \rho \quad \rho = \text{pore fluid density} \left[ \frac{M}{L^3} \right]$$

$$\phi = \text{porosity} \left[ \frac{L^3}{L^3} = 1 \right]$$

$u$  is fluid mass per unit volume of porous medium

Note: assume porous medium is saturated

2) Define the mass flux of pore fluid

$$\mathbf{j} = \rho \mathbf{q} = \rho \phi \mathbf{v} = u \mathbf{v} \quad \mathbf{q} = \text{vol. flux} \quad \frac{L^3}{L^2 T} = \frac{L}{T}$$

$$\mathbf{v} = \text{ave. inst. velocity} \quad \frac{L}{T}$$

$\Rightarrow$  advective flux

fluid mass carried/advected through the porous medium by groundwater flow

3) Vol. source:  $\hat{f} = \rho f$

Substitute

$$\frac{\partial}{\partial t}(\phi \rho) + \nabla \cdot (\rho \mathbf{q}) = \rho f$$

fluid mass balance  
in porous medium

Constitutive laws:

1) Darcy's law:  $\mathbf{q} = -K \nabla h$

2) Equation of state:  $\rho = \rho(h)$        $h = \frac{p}{\rho g}$

$\Rightarrow$  one equation for one unknown  $h$

Incompressible flow:

$$\rho = \text{const}$$

$$\phi = \phi(\underline{x}) \quad \text{does not change with time}$$

$$\frac{\partial}{\partial t}(\rho \phi) = 0$$

mass balance:  $\nabla \cdot \mathbf{q} = f$

Darcy's law:  $\mathbf{q} = -K \nabla h$

$\Rightarrow$

$$-\nabla \cdot [K \nabla h] = f$$

Poisson Equ