

Advective Heat Transport

So far we have considered heat conduction

$$\rho c_p \frac{\partial T}{\partial t} - \nabla \cdot \kappa \nabla T = \rho H$$

In lecture 2 we derived general balance law

$$\frac{\partial u}{\partial t} + \nabla \cdot \underline{j} = f_s$$

$u = \rho c_p (T - T_0)$ is unknown heat

$\underline{j} = -\kappa \nabla T = \underline{j}_D$ is conductive heat flux = \underline{j}_D

$f_s = \rho H$ is radiogenic heat production

Multiple fluxes and source terms

$$\frac{\partial u}{\partial t} + \nabla \cdot \sum_{i=1}^{N_f} \underline{j}_i = \sum_{i=1}^{N_s} f_i$$

$N_f = \#$ of fluxes

$N_s = \#$ of sources

Today add advective flux

general:

$$\underline{j}_A = \underline{v} u$$

$\underline{v} =$ velocity $[\frac{L}{T}]$

heat:

$$\underline{j}_A = \underline{v} \rho c_p (T - T_0)$$

$$\text{Heat flux} = \text{Thermal Energy Flux} = \frac{J}{L^2 T} = \frac{ML^2}{T^2} \frac{1}{L T} = \frac{M}{T^3}$$

$$\text{Energy: } J = Nm = \frac{ML}{T^2} L = \frac{ML^2}{T^2}$$

check units of advective heat flux

$$\underline{v} = \left[\frac{L}{T} \right] \quad \rho = \left[\frac{M}{L^3} \right] \quad c_p = \left[\frac{J}{kg K} = \frac{L^2}{T^2 \Theta} \right] \quad T = [\Theta]$$

$$\underline{j}_A = \frac{L}{T} \frac{M}{L^3} \frac{L^2}{T^2 \Theta} \Theta = \frac{M}{T^3} \checkmark$$

Adding advective flux to energy balance

$$\frac{\partial u}{\partial t} + \nabla \cdot [\underline{j}_A + \underline{j}_D] = f_s$$

$$\Rightarrow \rho c_p \frac{\partial T}{\partial t} + \nabla \cdot [\underline{v} \rho c_p (T - T_0) - \kappa \nabla T] = \rho H$$

$$\nabla \cdot [\underline{v} \rho c_p T_0] = \rho c_p T_0 \nabla \cdot \underline{v} = 0$$

Energy balance: $\rho c_p \frac{\partial T}{\partial t} + \nabla \cdot [\underline{v} \rho c_p T - \kappa \nabla T] = \rho H$

if $\kappa = \rho H = 0$ and $\rho c_p = \text{const.}$

$$\frac{\partial T}{\partial t} + \nabla \cdot [\underline{v} T] = 0$$

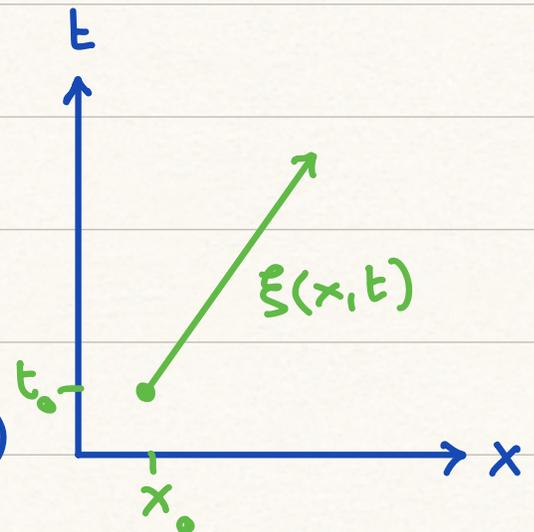
$$\frac{\partial T}{\partial t} + \underline{v} \cdot \nabla T = 0$$

Solving Advection with Method of Characteristics

Consider 1D heat transport

$$\text{PDE: } \frac{\partial T}{\partial t} + v \frac{\partial T}{\partial x} = 0 \quad x \in \mathbb{R}$$

$$\text{IC: } T(x, 0) = T_0(x)$$



Idea: Find a coordinate $\xi(x, t)$

along which PDE reduces

to an ODE. $T(x, t) \equiv \theta(\xi(x, t))$

$\xi \equiv$ characteristic curve

Total change of T along characteristic:

$$\frac{d\theta}{d\xi} = \frac{\partial T}{\partial t} \frac{dt}{d\xi} + \frac{\partial T}{\partial x} \frac{dx}{d\xi}$$

$$\text{PDE: } \frac{\partial T}{\partial t} + v \frac{\partial T}{\partial x} = 0$$

comparing

$$1) \frac{d\theta}{d\xi} = 0$$

$$2) \frac{dt}{d\xi} = 1$$

$$3) \frac{dx}{d\xi} = v \quad \left. \vphantom{\frac{dx}{d\xi} = v} \right\} \frac{dx}{dt} = v$$

Solve for char. eqn: $x - x_0 = v(t - t_0) \quad t_0 = 0$

$$t_0 = 0 : \quad x_0 = \underbrace{x - vt}$$

travelling wave coord.

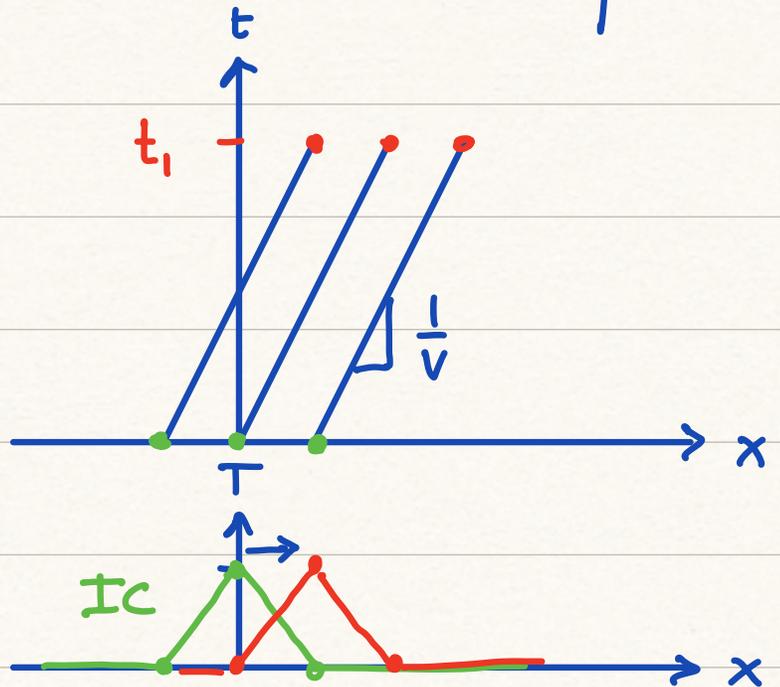
Substitute into IC

$$T(x, t=0) = T_0(x_0) = T_0(x-vt)$$

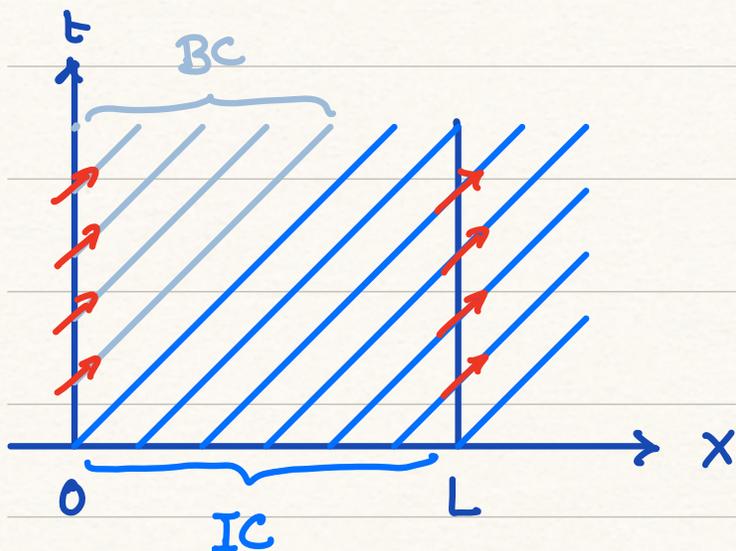
General Solution to linear Advection Equ

$$T(x, t) = T_0(x-vt)$$

IC is shifted to the right ($v > 0$) without change in shape! ∇
 \Rightarrow graphical interp.



Now consider a finite domain



- don't need outflow BC
 - need BC on inflow side
- \Rightarrow GEO 366M Math. Methods

\Rightarrow use in discretization

Steady Advection

Example: Adiabatic mantle upwelling

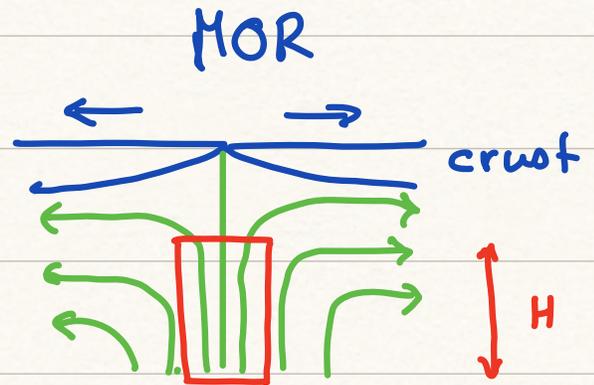
→ Mid-Ocean Ridge (MOR)

Corner flow in mantle

⇒ steady advection

Decompression cools mantle

⇒ sink term



$$\nabla \cdot [\underline{v} \rho c_p T] = -\alpha \rho g T \underline{v} \cdot \hat{z} \quad \rho c_p = \text{const}$$

$$\nabla \cdot [\underline{v} T] = -\frac{\alpha g T}{c_p} \underline{v} \cdot \hat{z} \quad \alpha = \text{thermal expansion coef.}$$
$$= 4 \cdot 10^{-5} \frac{1}{\text{K}}$$

(GEO 325C Cont. Mech.)

Directly beneath ridge \underline{v} is approx. vertical

$$\frac{d}{dz} [v T] = -\frac{\alpha g T v}{c_p} \quad \text{if } v = \text{const.}$$

$$\Rightarrow \frac{dT}{dz} = -\frac{\alpha g T}{c_p}$$

adiabatic T gradient

$$\alpha = 4 \cdot 10^{-5}$$

$$T = 1500 \text{ K}$$

$$\frac{dT}{dz} \approx 5 \cdot 10^4 \frac{\text{K}}{\text{m}}$$

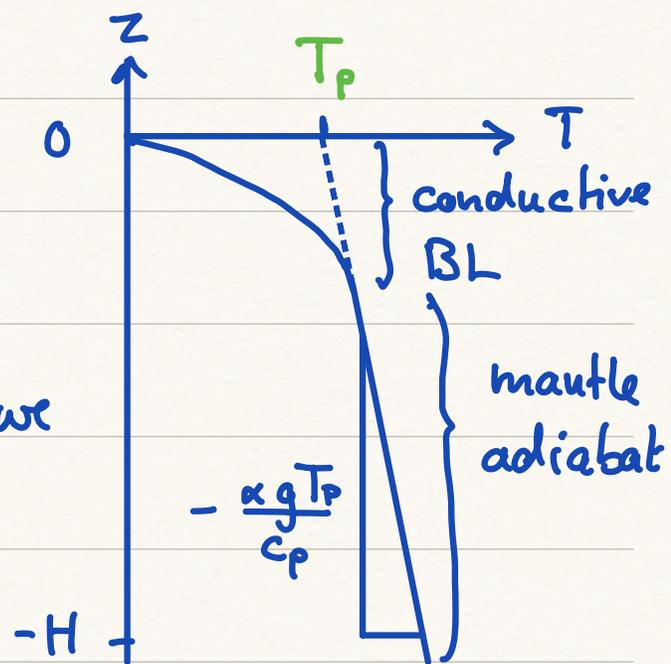
$$c_p = 1260 \frac{\text{J}}{\text{kg K}}$$

$$\approx 0.5 \frac{\text{K}}{\text{km}}$$

$$g = 9.81 \frac{\text{m}}{\text{s}^2}$$

To solve we need BC, but inflow bud is arbitrary

⇒ mantle potential temperature = adiabat at surface. = T_p



ODE: $\frac{dT}{dz} = -\frac{\alpha g}{c_p} T, \quad z \in [-H, 0]$

BC: $T(z=0) = T_p$

Integrate: $T = a \exp\left(-\frac{\alpha g}{c_p} z\right)$
 $T(z) = a e^0 = T_p$

Mantle adiabat:

$$T(z) = T_p e^{-\frac{\alpha g}{c_p} z}$$

$$\frac{\alpha g}{c_p} \approx \frac{5 \cdot 10^{-4} \cdot 10}{10^3} = 5 \cdot 10^{-6} \frac{1}{m} \ll 1$$

For $|z| \ll 10^6 \text{ m} = 10^3 \text{ km}$:

$$T(z) \approx -\frac{\alpha g T_p}{c_p} z$$

adiabatic gradient